

Phaselocking (and lack thereof) in a model for glacial variability

David Morawski

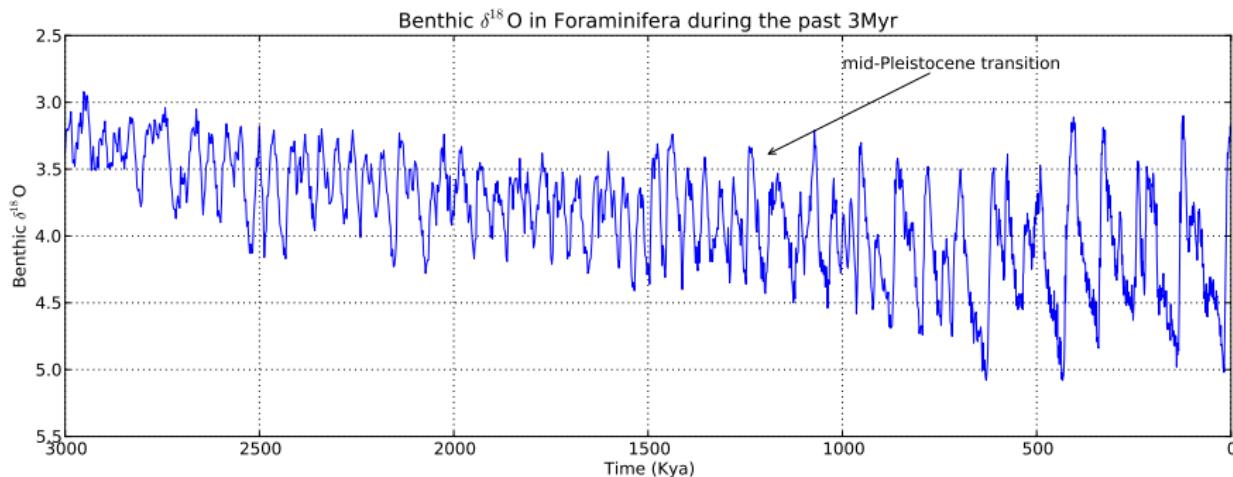
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Outline

- **Background:**
 - The mid-Pleistocene transition (**MPT**)
 - Huybers' explanation for the MPT
 - Huybers' model
- **Results:** Looking at Huybers' model
 - Constant ice accumulation
 - Randomizing ice accumulation
- **Huybers' versus Hopf**

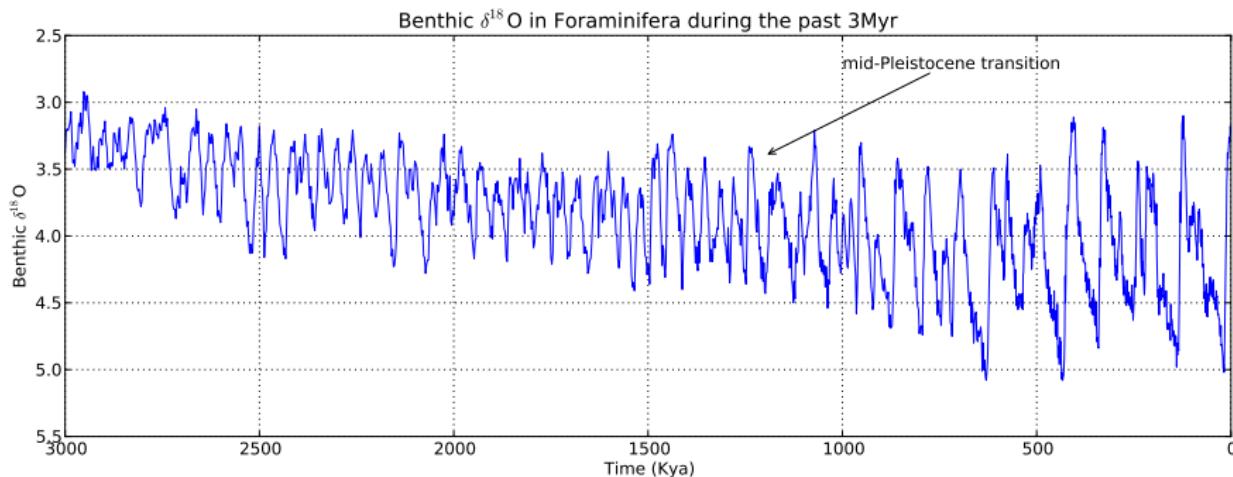
The mid-Pleistocene transition



- Before 1.2 Mya, glacial cycles had a period of 41Kyr.
- Since 1.2 Mya, the period has been 100Kyr.

Lisiecki, L.E. and M.E. Raymo. 2005. A Pliocene-Pleistocene stack of 57 globally distributed benthic D₁₈O records. *Paleoceanography*, Vol. 20, PA1003, doi:10.1029/2004PA001071.

The mid-Pleistocene transition

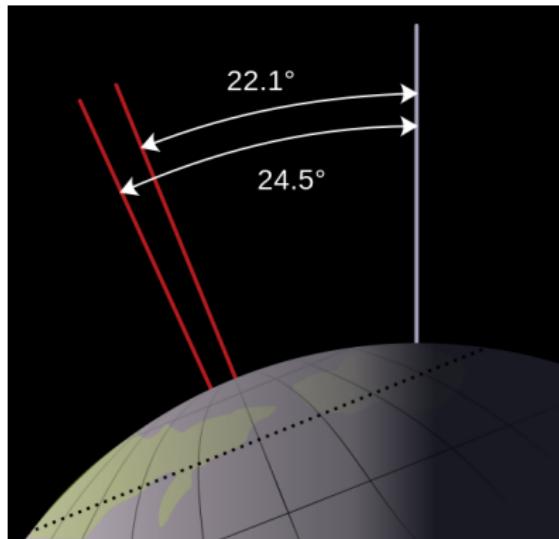


- Before 1.2 Mya, glacial cycles had a period of 41Kyr.
- Since 1.2 Mya, the period has been 100Kyr.
- “The 100,000 year problem”: what happened?

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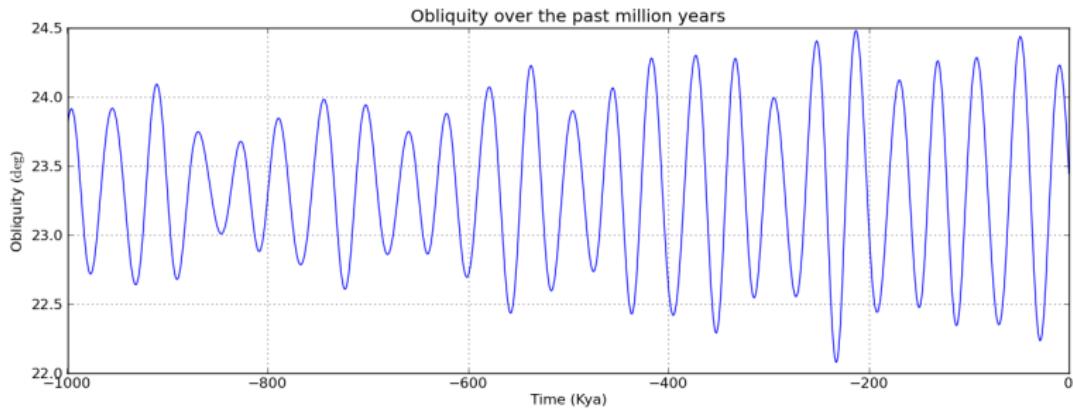
“Pacemaker of the Ice-Ages”

- Glacial cycles are driven by 3 orbital variations (Milankovitch cycles)
- The one that we care about today is **obliquity** (tilt):



Obliquity

Obliquity: 41Kyr period



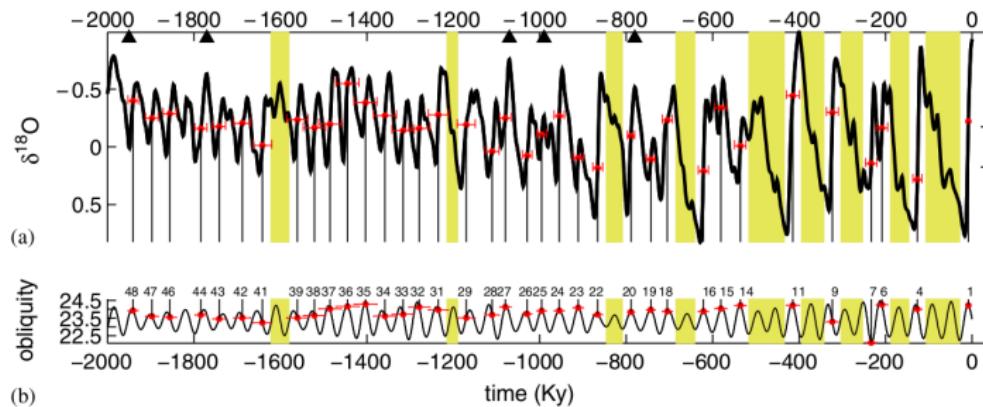
Berger A. and Loutre M.F., 1991. Insolation values for the climate of the last 10 million years. Quaternary Sciences Review, Vol. 10 No. 4, pp. 297-317, 1991.

Huybers' explanation of the mid-Pleistocene transition

- Obliquity triggered deglaciations throughout the Pleistocene
- Early Pleistocene: each period of obliquity triggered a deglaciation (41Kyr)
- Late Pleistocene: deglaciations skipped two or three obliquity cycles
 - (2×41) and (3×41) Kyr cycles averaged to 100Kyr

Huybers' explanation of the mid-Pleistocene transition

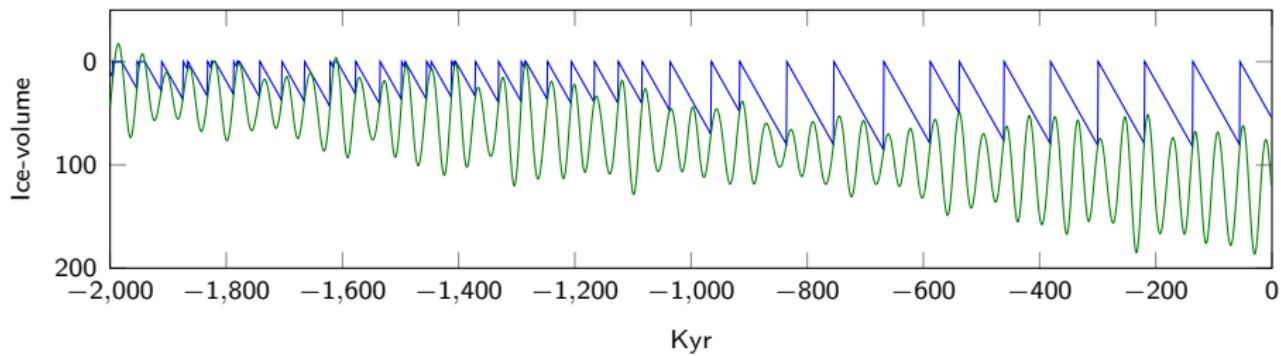
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Huybers P., 2007. Glacial variability over the last two million years: an extended depth-derived agemodel, continuous obliquity pacing, and the Pleistocene progression. *Quaternary Science Reviews* 26 (2007) 3755. ↗ ↘ ↙

Huybers' model

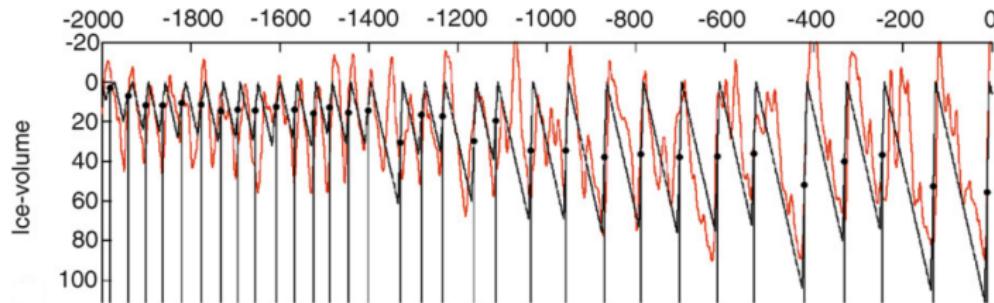
$$V_o = 10, a = 0.05, b = 126, c = 30$$



$$V_t = V_{t-1} + \eta_t \quad \text{and if } V_t \geq T_t \text{ terminate,}$$
$$T_t = at + b + c\theta'_t$$

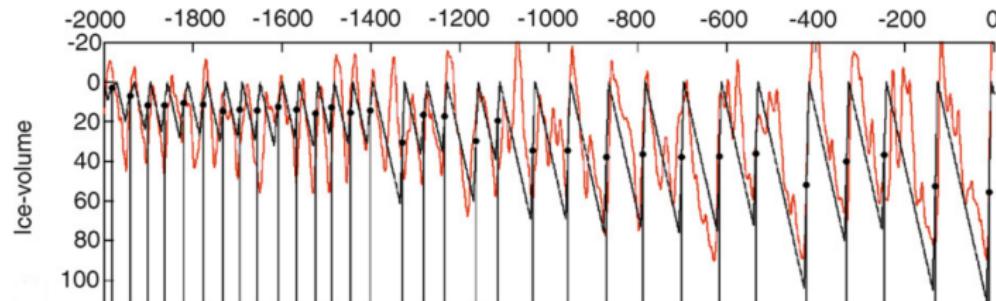
V_t :	ice volume
T_t :	threshold
η_t :	ice-accumulation
θ'_t :	obliquity (normalized)

Huybers' model



- “Selecting a slope of $a = 0.05\text{Ka}^{-1}$, an intercept of $b = 126$, and an obliquity amplitude of $c = 20$ reproduces the timing of most deglaciations over the last 2 Ma.”

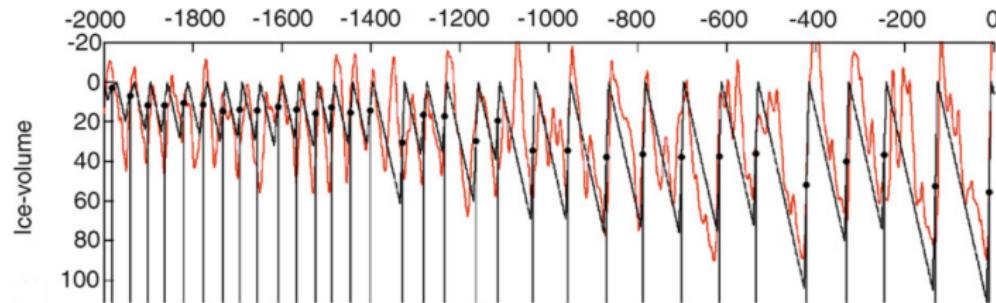
Huybers' model



- "Selecting a slope of $a = 0.05\text{Ka}^{-1}$, an intercept of $b = 126$, and an obliquity amplitude of $c = 20$ reproduces the timing of most deglaciations over the last 2 Ma."
- "Exceptions are that a deglaciation near 1.35 Ma is missed, the long glacial cycle at 1.6 Ma is not reproduced, . . . , and some of the smaller late-Pleistocene deglaciations."

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- “Exceptions are that a deglaciation near 1.35 Ma is missed, the long glacial cycle at 1.6 Ma is not reproduced, . . . , and some of the smaller late-Pleistocene deglaciations.”
- Also: the amplitudes of the early-Pleistocene are too small.

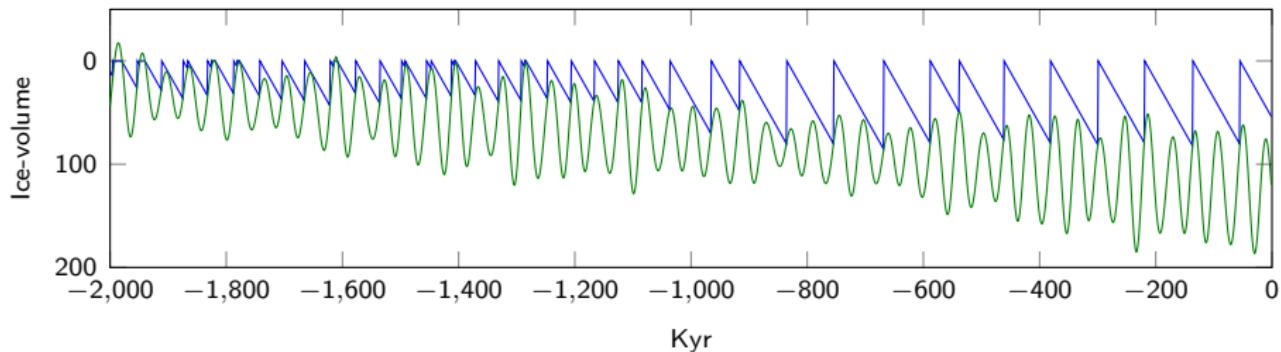
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How does the model behave when varying parameters and initial condition?

Huybers' model with constant ice accumulation

Set $\eta_t \equiv 1$:

$$V_o = 10, a = 0.05, b = 126, c = 30$$

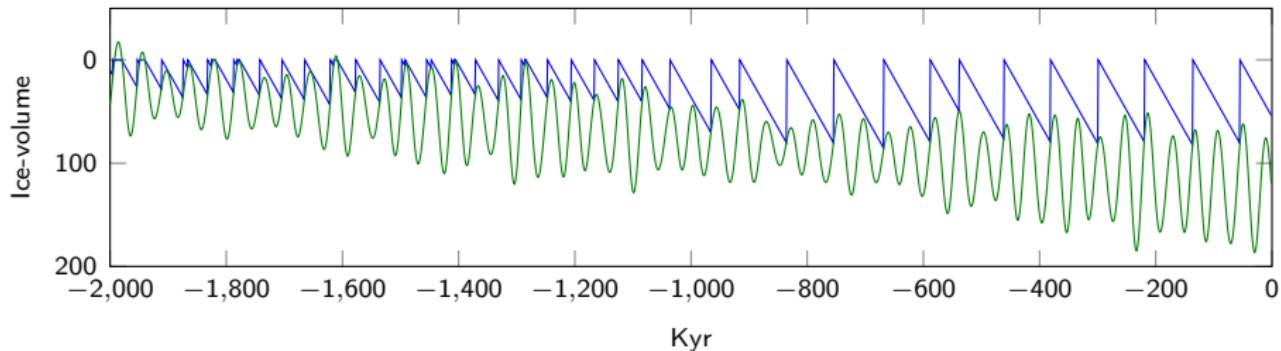


Q: For fixed b, c , how does initial ice volume v_o effect the model?

Huybers' model with constant ice accumulation

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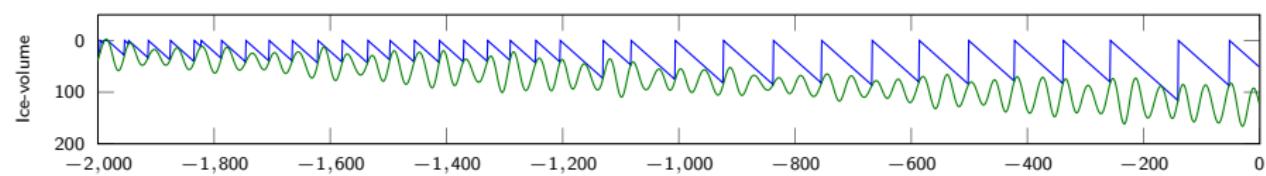
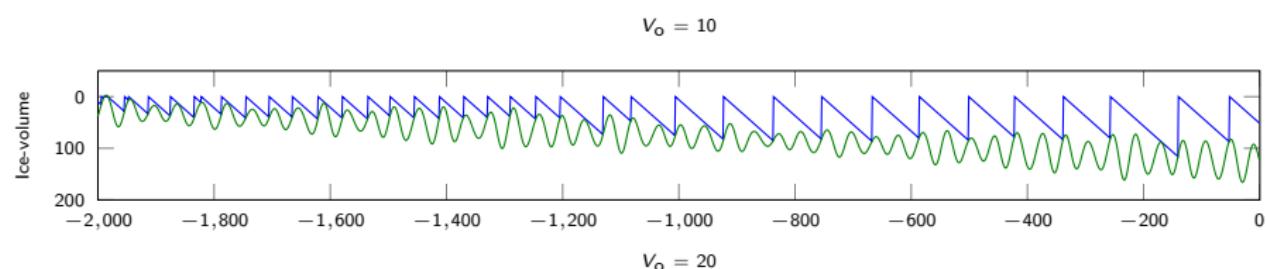
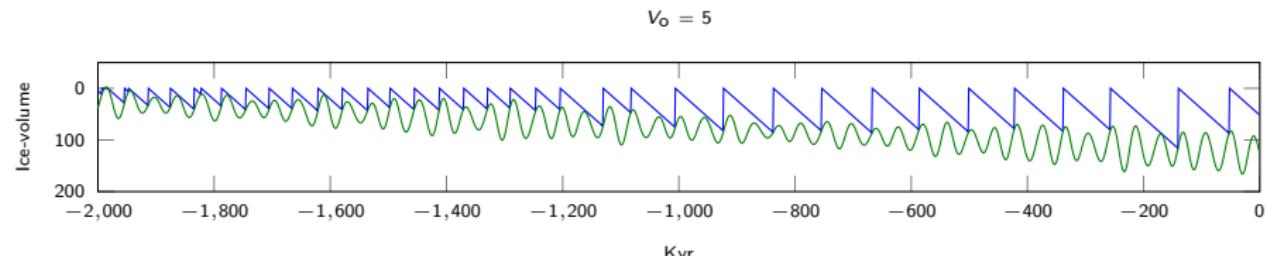


Q: For fixed b, c , how does initial ice volume V_o effect the model?

A: For most $b \in [0, 130]$ and $c \in [0, 30]$ fixed, (V_t, V_o) is eventually the same.

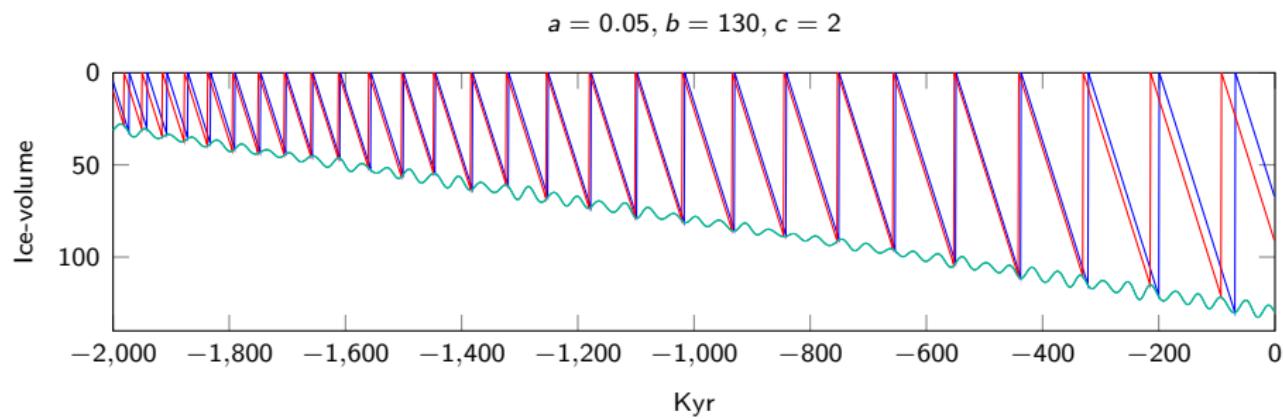
Huybers' model with constant ice accumulation

For most $b \in [0, 130]$ and $c \in [0, 30]$ fixed, (V_t, v_o) is eventually the same:



Huybers' model with constant ice accumulation

Sometimes, though, the phase is effected:

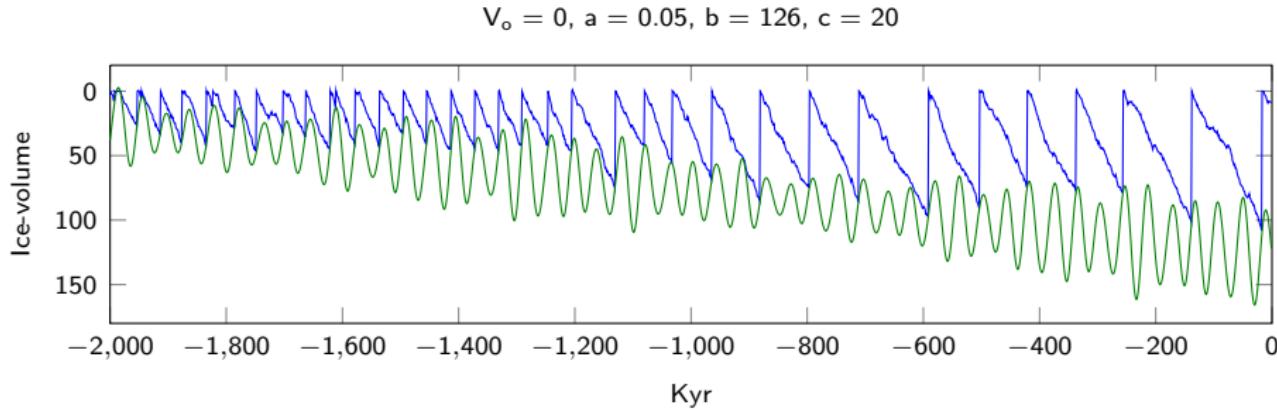


$$v_o = 5 \text{ versus } v_o = 10$$

Though this is mostly for small c .

Huybers' model with random ice accumulation

Now let η_t be a random variable ($\mu = \sigma = 1$)

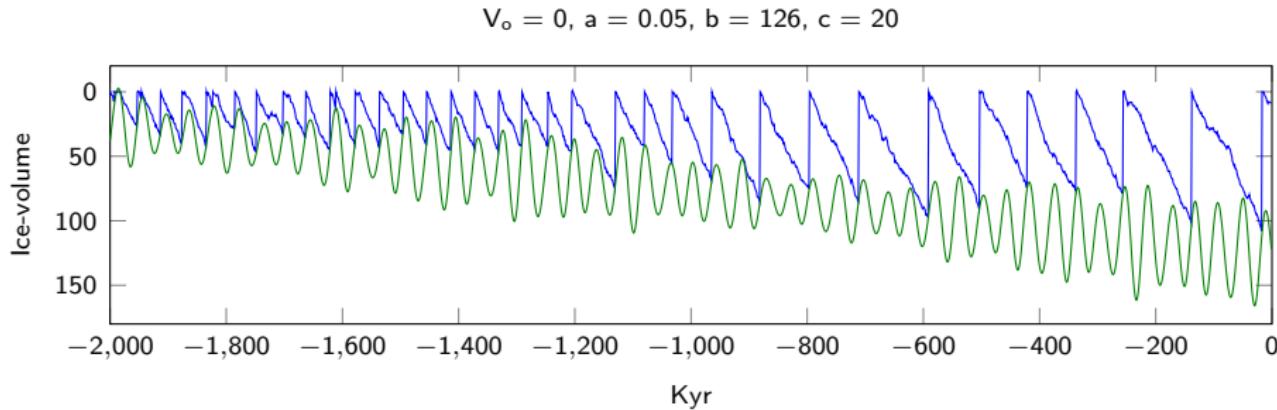


"The addition of a stochastic component to the model simulates the presence of weather at the highest frequencies and the myriad climatic processes not resolved by the model at longer periods."

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Huybers' model with random ice accumulation

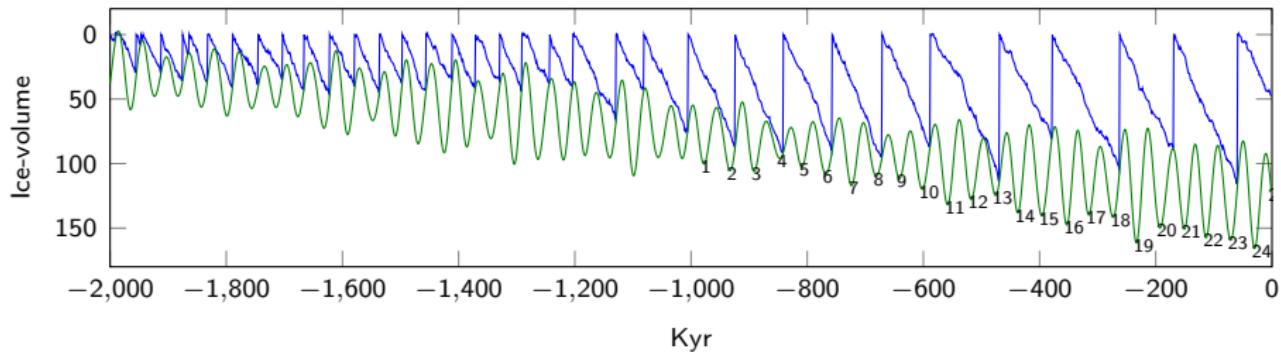
Now let η_t be a random variable ($\mu = \sigma = 1$)



"The timing of deglaciation is still controlled by obliquity, but obliquity cycle skipping is now random so that the glacial sequence need not coincide with the $\delta^{18}\text{O}$ stack."

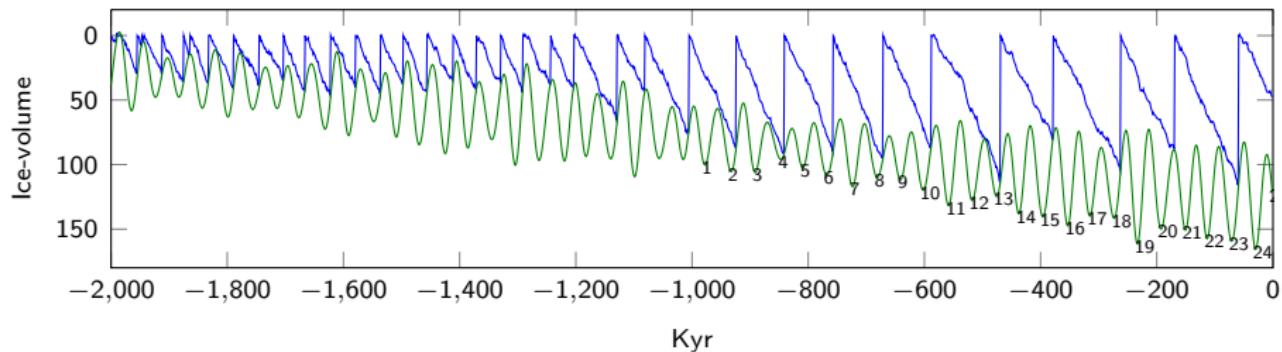
Random cycle skipping

$$V_0 = 0, a = 0.05, b = 126, c = 20, \sigma = 1$$

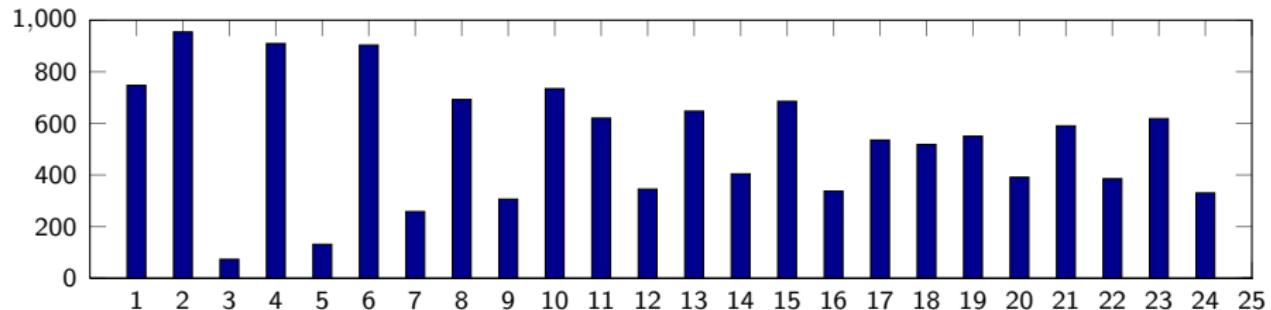


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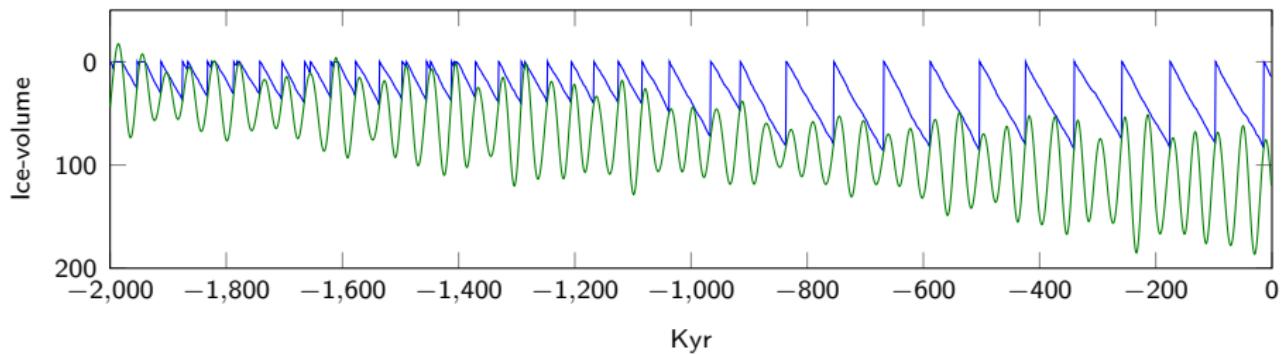


Trials = 1000

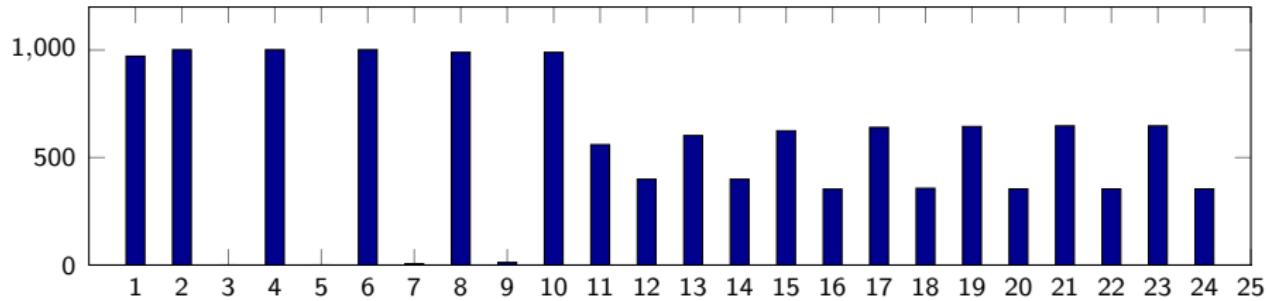


Random cycle skipping – small variance

$$V_0 = 0, a = 0.05, b = 126, c = 30, \sigma = 0.1$$

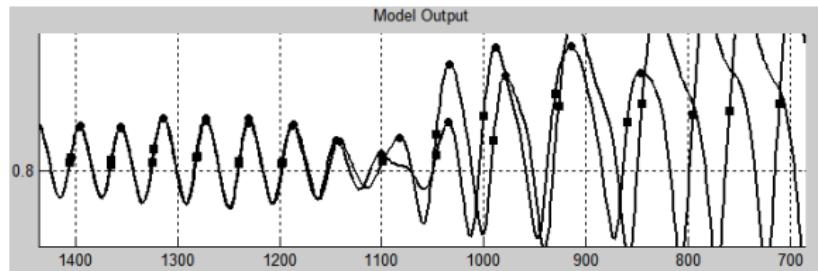


Trials = 1000



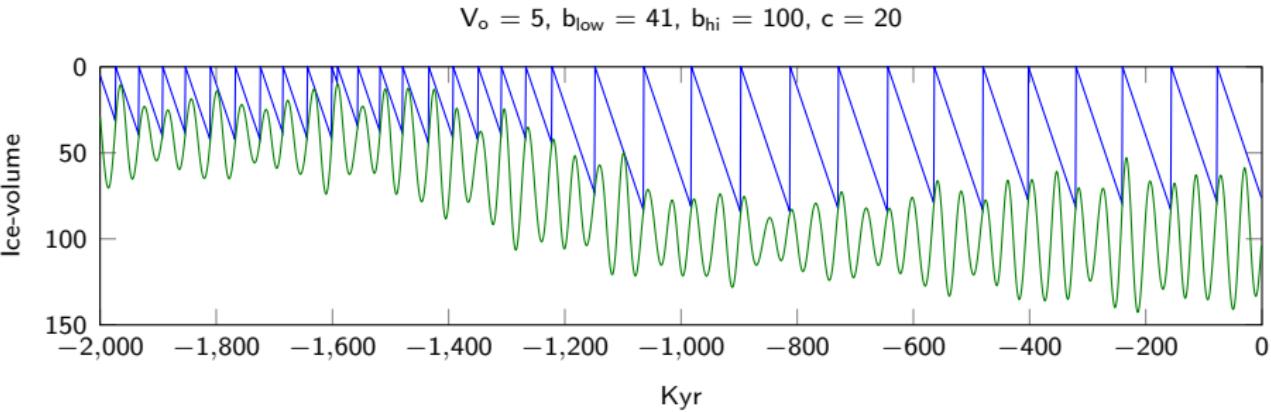
Huybers versus Hopf

Initial Conditions of Dynamic Hopf



One example of Sensitivity to Initial Condition at the neck.

An altered threshold function



$$T_t = \begin{cases} b_{\text{low}} + c\theta'_t & \text{if } t \in \{\text{early Pleistocene}\} \\ b_{\text{hi}} + c\theta'_t & \text{if } t \in \{\text{late Pleistocene}\} \end{cases}$$

An altered threshold function

In the deterministic case: very similar.

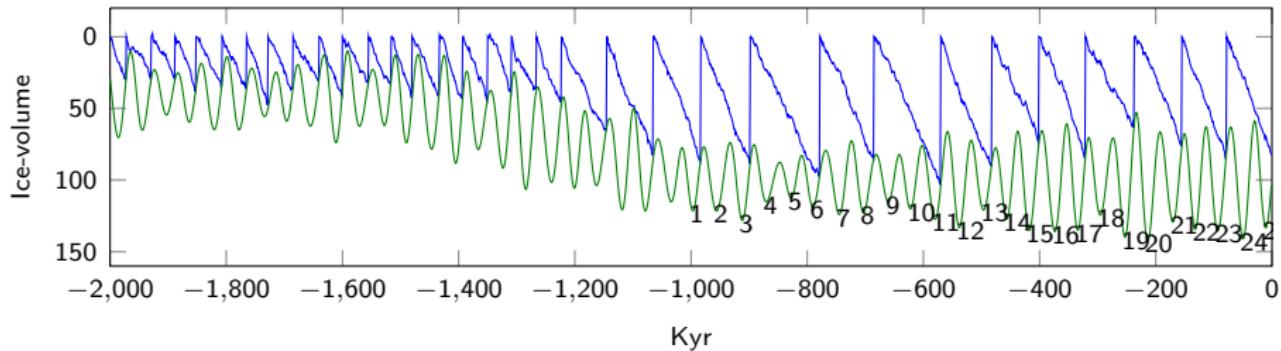
An altered threshold function

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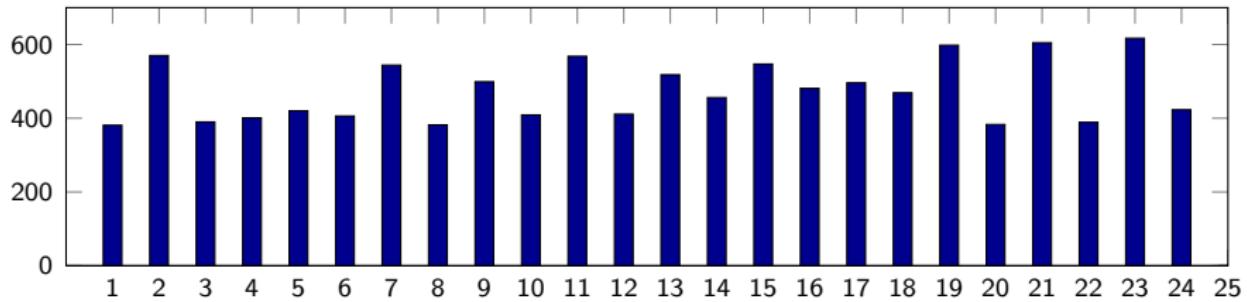
In the stochastic case: very similar.

Altered threshold: stochastic

$V_o = 0$, $b_{low} = 41$, $b_{hi} = 100$, $c = 20$ variance = 1

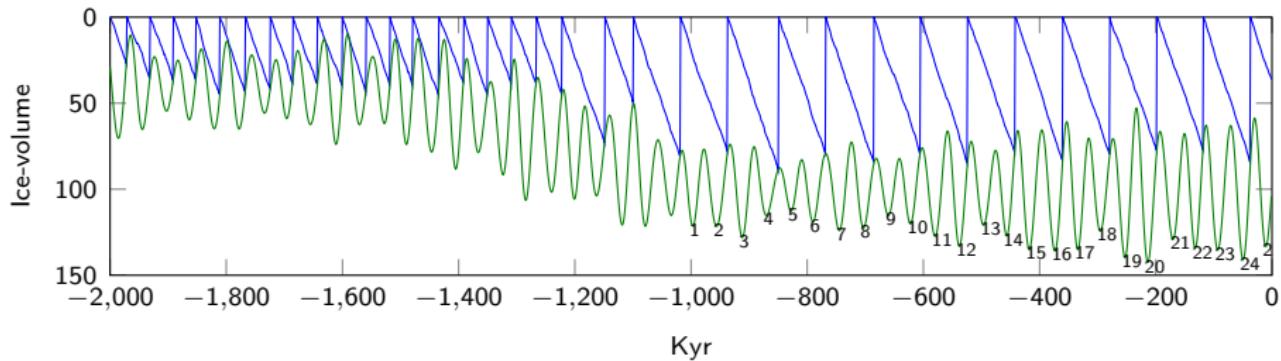


Trials = 1000, $\sigma = 1$

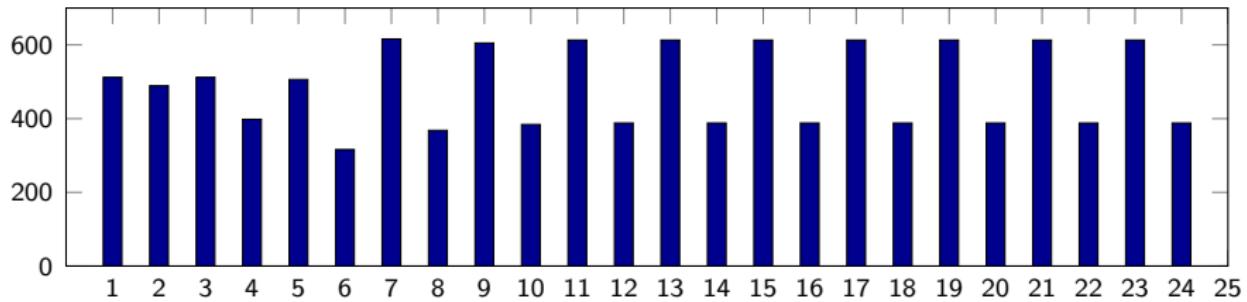


Altered threshold: stochastic, small variance

$$V_0 = 0, b_{\text{low}} = 41, b_{\text{hi}} = 100, c = 20, \sigma = 0.1$$



Trials = 1000



Thanks!